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N fluxes in two nitrogen saturated forested catchments in Germany: dynamics and modelling with INCA

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Abstract

The N cycle in forests of the temperate zone in Europe has been changed substantially by the impact of atmospheric N deposition. Here, the fluxes and concentrations of mineral N in throughfall, soil solution and runoff in two German catchments, receiving high N inputs are investigated to test the applicability of an Integrated Nitrogen Model for European Catchments (INCA) to small forested catchments. The Lehstenbach catchment (419 ha) is located in the German Fichtelgebirge (NO Bavaria, 690–871 m asl.) and is stocked with Norway spruce (*Picea abies* (L.) Karst.) of different ages. The Steinkreuz catchment (55 ha) with European beech (*Fagus sylvatica* L.) as the dominant tree species is located in the Steigerwald (NW Bavaria, 400–460 m asl.). The mean annual N fluxes with throughfall were slightly higher at the Lehstenbach (24.6 kg N ha⁻¹) than at the Steinkreuz (20.4 kg N ha⁻¹). In both catchments the N fluxes in the soil are dominated by NO₃. At Lehstenbach, the N output with seepage at 90 cm soil depth was similar to the N flux with throughfall. At Steinkreuz more than 50 % of the N deposited was retained in the upper soil horizons. In both catchments, the NO₃ fluxes with runoff were lower than those with seepage. The average annual NO₃ concentrations in runoff in both catchments were between 0.7 to 1.4 mg NO₃-N L⁻¹ and no temporal trend was observed. The N budgets at the catchment scale indicated similar amounts of N retention (Lehstenbach: 19 kg N ha⁻¹ yr⁻¹; Steinkreuz: 17 kg N ha⁻¹ yr⁻¹). The parameter settings of the INCA model were simplified to reduce the model complexity. In both catchments, the NO₃ concentrations and fluxes in runoff were matched well by the model. The seasonal patterns with lower NO₃ runoff concentrations in summer at the Lehstenbach catchment were replicated. INCA underestimated the increased NO₃ concentrations during short periods of rewetting in late autumn at the Steinkreuz catchment. The model will be a helpful tool for the calculation of “critical loads” for the N deposition in Central European forests including different hydrological regimes.

Keywords: forest ecosystem, modelling, N budgets, N saturation, NO₃ leaching, water quality, INCA

Introduction

The N cycle in many forest ecosystems of the temperate zone has been changed substantially by the impact of atmospheric N deposition. The rates of N deposition differ in various regions but, in Central Europe, 20–40 kg N ha⁻¹ yr⁻¹ are commonly found in throughfall (Dise *et al.*, 1998a, b). These inputs represent a significant disturbance of the N cycle. The effects of chronically high N deposition in forest ecosystems are often summarised as “N-saturation”. According to Aber *et al.* (1989), N saturation in forest ecosystems is reached when the deposition of N and N mineralisation exceed the biological N demand. This leads to NO₃ leaching into the groundwater and might increase the emission of N₂O and NO from the soils.

Nitrate concentrations in ground and surface water from forested catchments are normally far less than those from highly fertilised pastures and farmlands. However, even a small increase in NO₃ losses from forests may have large implications for drinking water management, since these waters are often used to dilute polluted water from other areas. Furthermore, NO₃ losses cause changes in the biogeochemistry of the catchments by increasing cation losses and inducing a higher risk of water acidification due to Al leaching (Dise *et al.*, 2001).

While substantial progress has been made in understanding the effects, predictions of the effects of N deposition on soil processes, trees, NO₃ leaching and denitrification in specific forest ecosystems are still

uncertain. At a regional scale, the NO_3 leaching from forests was related empirically to the N deposition and to the C/N ratio of the forest floor (Matzner and Grosholz, 1997; Dise *et al.*, 1998a, b). The positive influence of N deposition on the emission of N_2O and NO from forest soils remains a matter of debate (Brumme *et al.*, 1999).

Deterministic models of N turnover in terrestrial ecosystems have been developed for and applied mainly to agricultural systems. In these systems, the N cycle is driven by large inputs from fertilization and large outputs with biomass export, denitrification and seepage. (Whitehead, 1990; Jones *et al.*, 1991; Lunn *et al.*, 1996). In contrast, N inputs to forests from the atmosphere occur only at relatively low rates and are distributed more or less homogeneously throughout the year. The outputs by denitrification and seepage are generally much lower (Brumme *et al.*, 1999). Furthermore, the accumulation of N in biomass is small, while its immobilization in forest soils may be substantial (Magill *et al.*, 1997).

Here, data are presented on N fluxes and concentrations in two different forested catchments in Germany subjected to high rates of atmospheric N deposition. Long term developments and seasonal dynamics of NO_3 in runoff and on N budgets at the catchment scale are examined and the information collected is used to test the INCA model (Integrated Nitrogen Model for European Catchments; Whitehead *et al.*, 1998a, b) for its performance in predicting N fluxes for small forested catchments. The INCA model simulates flow pathways, N turnover including uptake and denitrification, fluxes and concentrations of both nitrate-N and ammonium-N in the soil, the groundwater and the river in daily resolution.

Sites and methods

SITES

The Steinkreuz catchment (55 ha in area) is located in the German Steigerwald area, $49^\circ 52' 20''$ N, $10^\circ 27' 40''$ E, at elevations between 400–460 m asl. The average air temperature is 7.5°C and mean annual precipitation is some 750 mm. Bedrock is Triassic sandstone (Middle Keuper) with clayey layers. Dominating soil types are Dystric Cambisols and Vertic Cambisols (FAO-System). Gleyic Cambisols occur only on small areas close to the creek. The soils are of sandy to loamy texture (Gerstberger, 2001). The forest stand comprises 75% of 130-year old European beech (*Fagus sylvatica* L.) with 25% sessile oak (*Quercus petraea* (Matt.) Liebl.) of the same age.

The Lehstenbach catchment (419 ha in area) is located in the German Fichtelgebirge area, $50^\circ 08' 35''$ N, $11^\circ 52' 10''$ E

at 690–871 m. The average air temperature is 5.8°C and the mean annual precipitation is 1100 mm. Bedrock is granite and the dominating soil types are acidic Cambisols, Gleysols and Histosols (FAO-System). A substantial area of the catchment is covered by wetlands and bogs (35%). The catchment is afforested by old Norway spruce (*Picea abies* (L.) Karst.) 45 to 160 years old. The stand at the plot 'Coulissenhieb' was planted in 1850. The Coulissenhieb site, located in the terrestrial part of the catchment with Cambisols and Cambic Podzols as dominating soil types (Gerstberger, 2001) has soil texture which varies from loamy sand to loam.

The C pools in the soil are higher at Coulissenhieb than at Steinkreuz (Table 1). The N pool of the Coulissenhieb soil is about 9400 kg ha^{-1} with 2900 kg ha^{-1} in the forest floor. At Steinkreuz, the total N pool of the soil was 6100 kg ha^{-1} with only 600 kg ha^{-1} in the forest floor.

Methods

At the Steinkreuz catchment, precipitation has been measured daily since January 1995. Samples for analysis of major elements have been collected bi-weekly with three replicates. Volume and element concentrations in throughfall have been measured bi-weekly in nine parallel samples from January 1995. Three of the nine samples are combined to give three mixed samples which are analysed for major elements. Stemflow amounts to about 7% of the soil water input and was measured at five trees. Samples of stemflow water were analysed for major elements bi-weekly. Ceramic suction lysimeters (seven replicates) were used to sample soil solution at 60 cm soil depth. Soil solutions have been collected and analysed bi-weekly since May 1995.

Discharge at the weir has been measured daily since November 1994 and water samples have been taken bi-weekly since then. Measurements of air temperature, air humidity, solar radiation, wind speed and wind velocity have been made since November 1994 at an open field station.

At the Lehstenbach catchment, bulk precipitation and throughfall as well as soil solution samples were sampled at the 'Coulissenhieb' plot. The volume of bulk precipitation has been measured daily since January 1994 and samples for analysis of macro-elements are collected bi-weekly with five replicates. Volume and element concentrations in throughfall have been measured bi-weekly in 20 parallel samples since June 1992. Ceramic suction lysimeters (20 replicates) have been used to sample soil solution at 90 cm soil depth. Soil solution samples have been collected bi-weekly since June 1993. Discharge at the weir has been measured daily since November 1986. Water samples have been taken bi-weekly since November 1987 for analysis of

Table 1a. Soil properties and soil N storage at the Steinkreuz catchment.

	Depth [cm]	pH (H ₂ O)	pH (CaCl ₂)	C [g kg ⁻¹]	N [g kg ⁻¹]	C-store [t ha ⁻¹]	N-store [t ha ⁻¹]
L	3-2	5.30	4.70	445.0	20.4		
Of	2-0.5	5.10	4.50	413.0	21.8	12.0 ^a	0.6 ^a
Oh	0.5-0	4.10	3.40	205.0	11.5		
I Ah	0-5	3.90	3.20	66.9	4.3	35.5	2.3
I Bv	5-24	4.30	3.80	10.9	0.6	27.1	1.5
I Swd-Bv1	24-50	4.60	3.90	3.5	0.2	7.1	0.4
I Swd-Bv2	50-80	4.90	4.00	1.6	0.2	4.1	0.5
II Cv	80-85	5.20	4.10	1.4	0.3	0.6	0.1
III Cv	85-115	5.20	4.10	1.4	0.2	1.9	0.3
IV Cv1	>115	5.50	4.20	0.7	0.1	1.4	0.2
IV Cv2		5.50	4.20	0.8	0.1	1.3	0.2

^a = Sum of: L+Of+Oh

Table 1b. Soil properties and soil N storage at the Lehstenbach catchment.

	Depth [cm]	pH (H ₂ O)	pH (CaCl ₂)	TOC [g kg ⁻¹]	TON [g kg ⁻¹]	C-store [t ha ⁻¹]	N-store [t ha ⁻¹]
L	8.5-8	4.50	3.60	478.0	19.3	3.6	0.2
Of	8-3	3.80	2.90	372.0	18.0	25.4	1.2
Oh	3-0	3.50	2.60	376.0	16.6	31.0	1.5
Ahe	0-10	3.70	2.90	38.9	1.7	26.4	1.2
Bh	10-12	3.80	3.30	90.5	4.0	8.2	0.4
Bhs	12-30	4.40	3.90	53.6	3.8	49.3	3.5
BvCv	30-55	4.50	4.30	8.4	0.5	20.0	1.2
Cv1	55-70	4.50	4.20	2.2	0.2	1.5	0.1
Cv2	>70	4.50	4.10	2.0	0.2	0.9	0.1

major elements. Meteorological measurements of air temperature, air humidity, solar radiation, wind speed and wind velocity have been made since May 1994 at an open field station.

The water fluxes with seepage at the plot-scale were calculated by the soil water model Hydrus-2D (Simunek *et al.*, 1996) for Steinkreuz at 60 cm depth and by the model SIMULA18, (Manderscheid, 1995) at 90 cm depth for the Coulissenhieb plot. Both models simulate seepage based on estimates of actual transpiration, soil physical parameters and root distribution.

Parameterisation of the INCA model

The driving data for the INCA model, including soil moisture deficit, hydrologically effective rainfall, air

temperature, actual throughfall, discharge and inorganic nitrogen concentrations were compiled for the period 1995–1999. Soil water deficit was calculated as the difference between actual transpiration and throughfall. At the Steinkreuz, the calculated seepage at 60 cm depth was taken as hydrologically effective rainfall while at Lehstenbach, the hydrologically effective rainfall was calculated as throughfall – actual transpiration – soil water deficit.

At the Steinkreuz catchment, a discharge to deeper groundwater that is not accounted for at the weir measurements, is suggested by the fact that the CI budget (throughfall + stemflow – runoff) of the catchment is strongly positive over the five years using measured discharge rates. In addition, transpiration measurements at the trees revealed annual transpiration rates of only 180 mm (Tenhunen *et al.*, 2001) which supports the suggestion of

Table 2. Parameters of the INCA model used for the modelling.

	<i>Steinkreuz</i>	<i>Lehstenbach</i>
Denitrification m day^{-1}	0.002	0.007
Nitrogen fixation ($\text{kg N ha}^{-1} \text{ day}$)	0	0
Nitrification m day^{-1}	0.02	0.025
Mineralisation ($\text{kg N ha}^{-1} \text{ day}$)	0	0
Immobilisation m day^{-1}	0	0
Soil moisture deficit maximum (mm)	40	40
Max. temperature difference ($^{\circ}\text{C}$)	8	4.5
Sustain Surface flow at ($\text{m}^3 \text{ s}^{-1}$)	0	0
Sustain Sub-surface flow at ($\text{m}^3 \text{ s}^{-1}$)	0	0.007
Stop denitrification at ($^{\circ}\text{C}$)	0	0
Stop nitrification at ($^{\circ}\text{C}$)	0	0
Fertilizer addition start day	0	0
Fertilizer addition period (days)	0	0
Nitrate addition rate ($\text{kg N ha}^{-1} \text{ day}^{-1}$)	0	0
Ammonium addition rate ($\text{kg N ha}^{-1} \text{ day}^{-1}$)	0	0
Plant growth start day	60	90
Plant growth period (days)	210	160
Nitrate uptake rate m day^{-1}	0.015	0.01
Ammonium uptake rate m day^{-1}	0.015	0.8
Maximum uptake ($\text{kg N ha}^{-1} \text{ yr}^{-1}$)	70	70
V_{rMax} (depth \times porosity)	0.1	0.2
Soil reactive zone (days)	1	1
Groundwater zone (days)	40	40
Initial Surface flow ($\text{m}^3 \text{ s}^{-1}$)	0.001	0
Initial Surface nitrate (mg l^{-1})	0.6	0.5
Initial Surface Drainage volume (m^3)	100	1000
Initial Sub-surface flow ($\text{m}^3 \text{ s}^{-1}$)	0.001	0
Initial Sub-surface nitrate (mg l^{-1})	1.3	0.99
Initial Sub-surface Drainage volume (m^3)	1E6	8E6
Instream Denitrification rate (day^{-1})	0.001	0.005
Instream Ammonium nitrification rate (day^{-1})	0	0.5
Reach parameter a	0.04	0.04
Reach parameter b	0.67	0.67
Base flow index	0.55	0.55

groundwater losses from the catchment. Thus, the discharge was adjusted to a balanced long term CI budget which resulted in an average discharge rate of 257 mm. The latter was also used to calculate the annual fluxes of elements in runoff.

The parameterisation of the model was simplified to reduce the complexity. The soil N pools were assumed to be in steady state (Table 2) and, subsequently, the rates of NH_4 mineralisation, NH_4 immobilisation and the NO_3 fixation were set to zero. Only one land-use type was postulated. Plant uptake was adjusted to give the N demand for net increment of the trees. To account for the larger area of hydromorphic sites, the denitrification rate for the Lehstenbach was set higher than for the Steinkreuz catchment. The greater size of the Lehstenbach leads to higher initial surface and sub-surface drainage volumes. Because of the higher elevation, the plant growth period at Lehstenbach is shorter than at Steinkreuz.

The NH_4 uptake is assumed to be higher in spruce stands than in beech stands (Berger, 1995).

Results

HYDROLOGY

The distribution of throughfall is rather homogenous throughout the year in both catchments (Fig. 1). Maximum

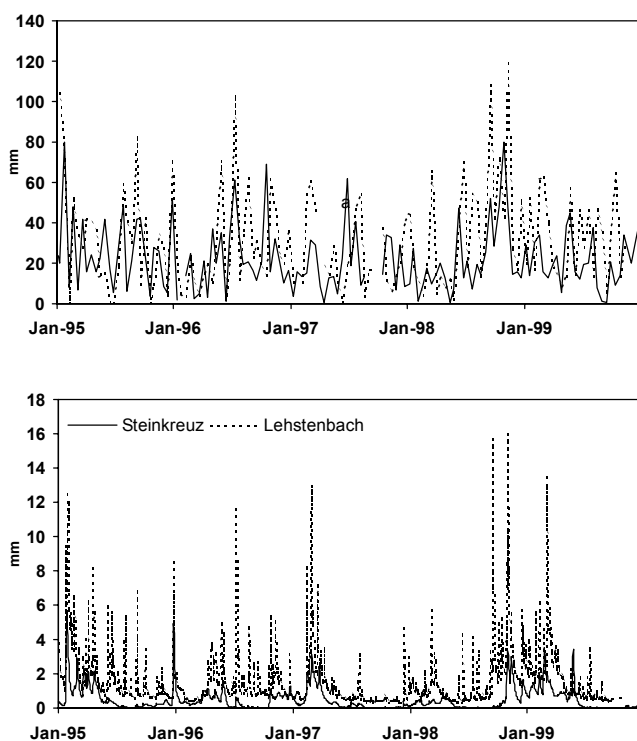
**Fig. 1.** Throughfall (biweekly) and runoff (daily) at both catchments

Table 3. Annual water fluxes with throughfall and discharge (in mm) at both catchments.

Year	Steinkreuz			Lehstenbach	
	TF+SF	Runoff ^a	Runoff ^b	TF	Runoff ^b
1995	723.8	251.4	318.8	836.1	689.4
1996	576.2	156.3	243.5	690.5	490.0
1997	483.7	132.0	183.5	645.5	416.6
1998	615.3	149.0	301.4	951.5	614.8
1999	589.5	191.5	239.4	745.3 ^c	482.5 ^c
Mean:	597.7	176.0	257.3	773.8	538.7

^a = Measured at the weir^b = Runoff according to CI-budget^c = 01/99-10/99: observed values; 11/99-12/99: estimated from long term average

TF = Throughfall

SF = Stemflow

rates of throughfall observed at the bi-weekly scale were about 80 mm at Steinkreuz and 120 mm at Lehstenbach. Runoff is at a minimum in summer in both catchments. Maximum daily discharge rates were 6 mm at Steinkreuz and 16 mm at Lehstenbach. The creek at Steinkreuz is almost dry during the late summer while, at the Lehstenbach creek, there was a baseflow of about 0.25 mm d⁻¹ in summer. At the Lehstenbach, snow cover and snow melt have a substantial effect on winter and spring hydrology, which is not the case for the Steinkreuz. The mean discharge from 1995–1999 was 539 mm at Lehstenbach, while in Steinkreuz only 176 mm (Table 3) was measured at the weir. The annual discharge rates from 1995 to 1999 varied by about 270 mm at the Lehstenbach but only by about 120 mm at Steinkreuz.

N-CONCENTRATIONS

Mineral N concentrations in throughfall were similar in both catchments (Fig. 2a, b and Table 5a, b). Furthermore, the concentrations of NH₄ and NO₃ in throughfall were almost equal. The concentrations of dissolved organic nitrogen (DON) in throughfall were substantial and reached about 29.2% of the mineral N at the Steinkreuz and 15.4% at the Lehstenbach. In contrast to throughfall, mineral N in soil solutions and runoff is almost completely represented by NO₃. The concentrations of DON in soil solutions and runoff are generally low and close to detection (< 0.2 mg N l⁻¹).

Concentrations of N in throughfall are highly variable with time while those in soil solutions had less short term variation. Soil solution concentrations of NO₃ showed a large spatial variation and some long term patterns in both plots: at Steinkreuz, the concentrations were generally low in 1998

and 1999 while at Lehstenbach, low concentrations were found in 1995 and 1999. The concentrations observed in runoff in both catchments are lower than average soil solution concentrations. In both catchments, the concentrations in runoff were similar with annual means between 0.7 and 1.4 mg NO₃-N l⁻¹. The temporal patterns found in the soil solutions were not reflected in runoff. In

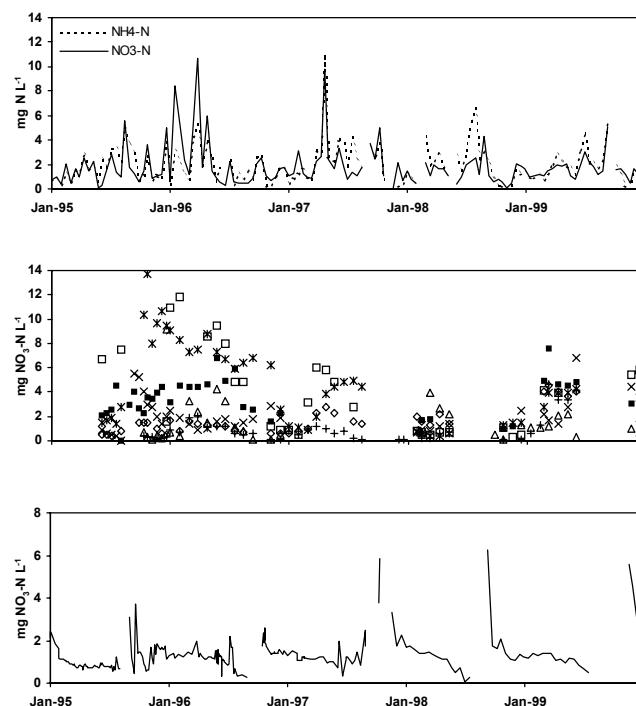


Fig. 2a. Nitrogen concentrations in throughfall, soil solutions and runoff at the Steinkreuz catchment

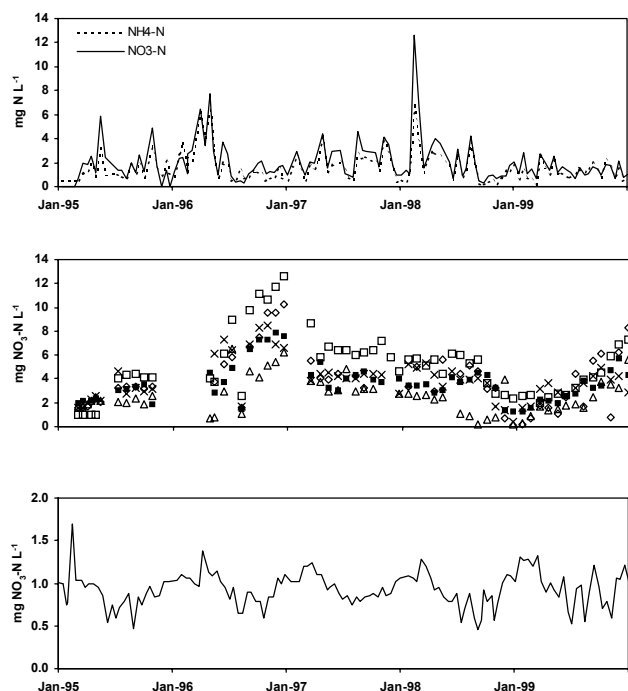


Fig. 2b. Nitrogen concentrations in throughfall, soil solutions and runoff at the Lehstenbach catchment

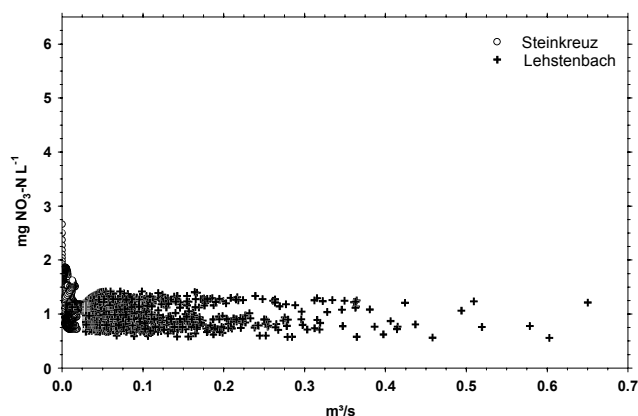


Fig. 3. Simulated nitrate concentrations in runoff as related to discharge

the runoff from the Lehstenbach, there was a seasonal pattern with lower concentrations in summer and autumn followed by higher concentrations in winter and spring. At Steinkreuz, peaks of NO_3 concentrations in runoff occurred shortly after rewetting in late autumn. No long term trend of NO_3 in runoff was observed in either catchment and the concentrations of NO_3 were not related to discharge volume (Fig. 3).

N FLUXES

The fluxes of mineral N with throughfall were on average higher at Lehstenbach ($21 \text{ kg N ha}^{-1} \text{ yr}^{-1}$) than at Steinkreuz ($15.7 \text{ kg N ha}^{-1} \text{ yr}^{-1}$; Table 4a, b). The fluxes of DON in

throughfall were remarkable with slightly higher fluxes at Steinkreuz. The annual variation in throughfall fluxes was low except for DON at Steinkreuz where fluxes varied between 3.7 to $5.6 \text{ kg N ha}^{-1} \text{ yr}^{-1}$.

The N fluxes with soil solution were clearly dominated by NO_3 in both catchments, but the fluxes differed between the catchments. At Lehstenbach, the seepage output of about $21 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ was almost equal to the $\text{NH}_4 + \text{NO}_3$ flux in throughfall, while at Steinkreuz seepage was about $6 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ (Table 4a). The annual variation of the NO_3 fluxes with seepage was high at Steinkreuz with a range of 3.8 to $9.4 \text{ kg N ha}^{-1} \text{ yr}^{-1}$. In contrast, the annual variation at the Coulissenhieb plot was low.

The NO_3 fluxes with runoff from both catchments had only small annual variations and no trend was visible. Fluxes of DON in seepage and runoff were negligible. On average, the runoff flux was about $4.9 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ at Lehstenbach and $3.3 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ at Steinkreuz. Thus, in both catchments the seepage flux of NO_3 was higher than with runoff. This observation was most pronounced at the Lehstenbach.

Taking the throughfall fluxes of N (including DON) as deposition input and runoff as output, both catchments retained (including gaseous emissions) large amounts of N. In case of the Steinkreuz, a retention of $17 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ was found on average (Table 4a). At Lehstenbach, the N retention was about $19 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ (Table 4b). The N budgets for the plots, calculated as throughfall-N (incl. DON) minus seepage, revealed differences: The average N retention at the Steinkreuz plot was about $14 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ which is similar to the retention at the catchment scale while only $3 \text{ kg N ha}^{-1} \text{ yr}^{-1}$ were retained at the Coulissenhieb plot which is far less than the retention observed at the catchment scale.

INCA APPLICATION

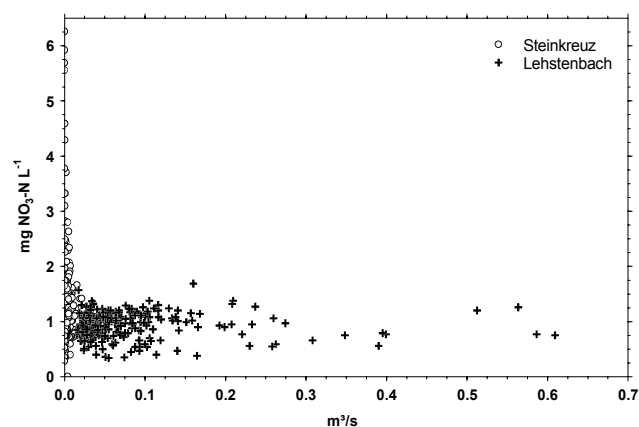


Fig. 4. Observed nitrate concentrations in runoff as related to discharge

Table 4a. Annual nitrogen fluxes ($\text{kg N ha}^{-1} \text{ yr}^{-1}$) with throughfall (TF), seepage in 60 cm soil depth (SP) and runoff (R) in the Steinkreuz catchment.

		1995	1996	1997	1998	1999	Mean
TF	$\text{NO}_3\text{-N}$	7.33	7.30	7.25	6.46	7.40	7.15
	$\text{NH}_4\text{-N}$	10.31	7.45	8.80	8.46	8.09	8.62
	DON	-	4.01	5.58	5.10	3.74	4.61
SP	$\text{NO}_3\text{-N}$	-	7.45	3.87	3.88	9.36	6.14
	$\text{NH}_4\text{-N}$	-	0.45	0.09	0.11	0.14	0.20
	DON	-	n.d.	n.d.	n.d.	n.d.	n.d.
R ^a	$\text{NO}_3\text{-N}$	3.39	3.89	2.28	3.95	3.28	3.36
	$\text{NH}_4\text{-N}$	0.06	0.11	0.04	0.06	0.05	0.06
	DON	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.

^a = Runoff volume adjusted to Cl-budget (see Table 3)

n.d. = not detectable

- = not measured

Table 4b. Annual nitrogen fluxes ($\text{kg N ha}^{-1} \text{ yr}^{-1}$) with throughfall (TF), seepage in 90 cm soil depth (SP) and runoff (R) in the Lehstenbach catchment.

		1995	1996	1997	1998	1999 ^a	Mean
TF	$\text{NO}_3\text{-N}$	14.85	10.89	12.01	12.73	9.76	12.05
	$\text{NH}_4\text{-N}$	11.02	9.21	9.25	9.07	7.87	9.28
	DON	2.23	3.02	4.26	3.68	3.23	3.28
SP	$\text{NO}_3\text{-N}$	21.36	24.49	20.03	20.25	18.09	20.84
	$\text{NH}_4\text{-N}$	0.39	0.49	0.28	0.30	0.30	0.35
	DON	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.
R	$\text{NO}_3\text{-N}$	5.94	5.20	4.14	4.27	5.09	4.93
	$\text{NH}_4\text{-N}$	0.21	0.29	0.17	0.22	0.23	0.22
	DON	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.

^a = 01/99-10/99: observed values; 11/99-12/99: estimated

Figure 3 shows the relation of the simulated daily $\text{NO}_3\text{-N}$ concentration to the simulated daily discharge. As is observed in practice (Fig. 4) the simulated runoff NO_3 concentrations for both catchments were not related to the volume of discharge.

The simulated discharge for the *Steinkreuz catchment* was much higher than that observed since the seepage water fluxes at the catchment scale were input as hydrologically

effective rainfall. Thus, the simulated discharge in INCA did not correspond well with the measurements (Fig. 5a).

The observed NO_3 concentrations in runoff in the Steinkreuz were generally matched well by the model (Fig. 5a). The measured annual average concentration of $\text{NO}_3\text{-N}$ is 1.32 mg l^{-1} while the model calculated an average concentration of 1.12 mg l^{-1} ; INCA underestimated the NO_3 concentration during rewetting after drought in late autumn.

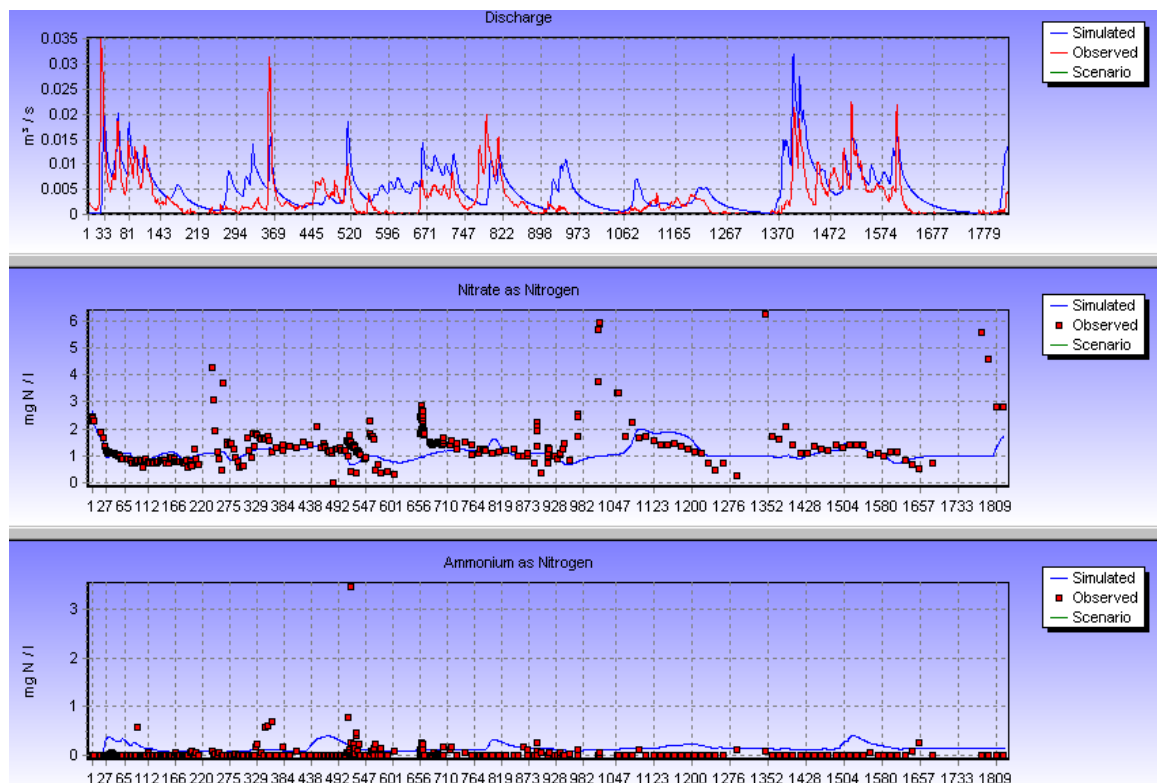


Fig. 5a. Simulated discharge and nitrogen concentrations in runoff of the Steinkreuz catchment

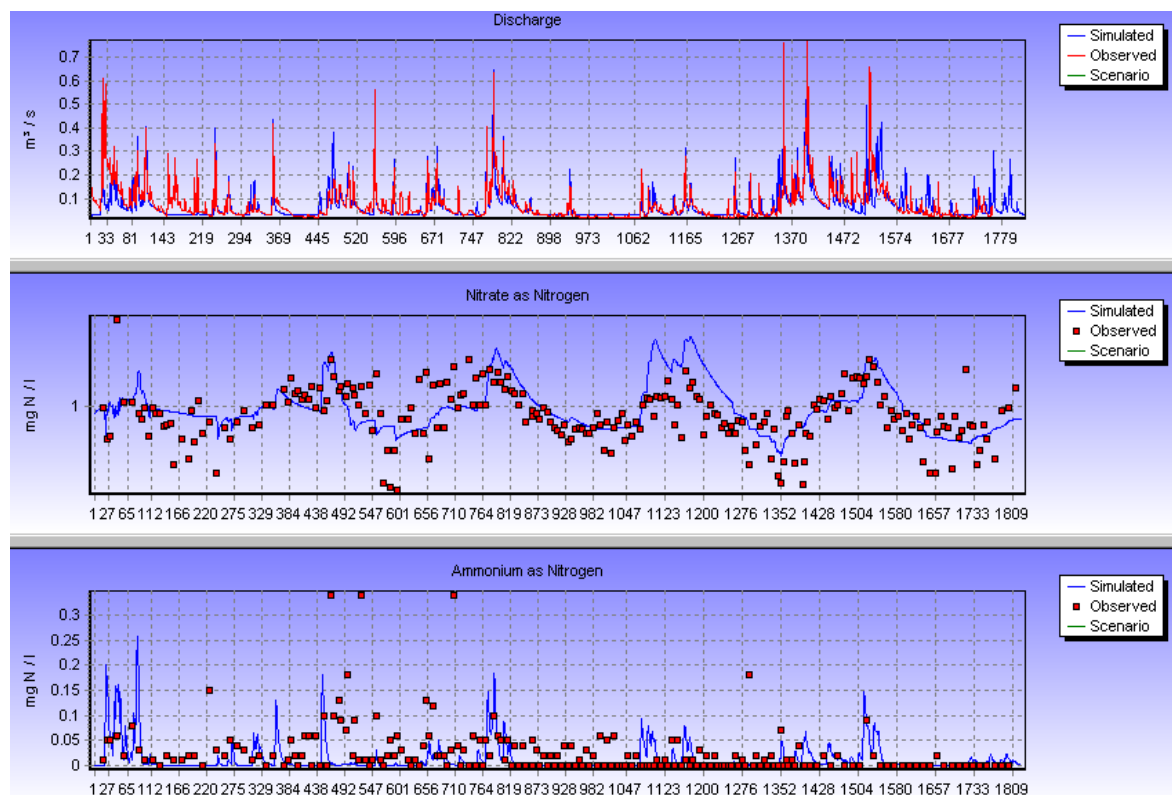


Fig. 5b. Simulated discharge and nitrogen concentrations in runoff of the Lehstenbach catchment

Table 5a. Average concentrations of nitrogen (mg N l⁻¹) in solutions of the Steinkreuz catchment

Year	TF	NO ₃ -N 60 cm depth	runoff	TF	60 cm depth	NH ₄ -N runoff	TF	DON 60 cm depth	runoff
1995	1.01		1.06	1.42		0.02		n.d.	n.d.
1996	1.27	2.42	1.60	1.29	0.15	0.04	0.70	n.d.	n.d.
1997	1.50	1.89	1.24	1.82	0.04	0.02	1.15	n.d.	n.d.
1998	1.05	1.24	1.31	1.37	0.04	0.01	0.83	n.d.	n.d.
1999	1.26	3.43	1.37	1.37	0.05	0.01	0.63	n.d.	n.d.
Mean:	1.22	2.25	1.32	1.45	0.07	0.02	0.83	n.d.	n.d.

Table 5b. Average concentrations of nitrogen (mg N l⁻¹) in solutions of the Lehstenbach catchment

Year	TF	NO ₃ -N 90 cm depth	runoff	TF	90 cm depth	NH ₄ -N runoff	TF	DON 90 cm depth	runoff
1995	1.78	4.15	0.86	1.32	0.08	0.03	0.27	n.d.	n.d.
1996	1.58	5.02	1.06	1.33	0.10	0.06	0.44	n.d.	n.d.
1997	1.86	4.48	0.99	1.43	0.06	0.04	0.66	n.d.	n.d.
1998	1.34	3.82	0.69	0.95	0.06	0.04	0.39	n.d.	n.d.
1999 ^a	1.31	3.74	1.05	1.06	0.06	0.05	0.43	n.d.	n.d.
Mean:	1.57	4.24	0.93	1.22	0.07	0.04	0.44	n.d.	n.d.

^a = 01/99-10/99: observed values; 11/99-12/99: estimated

In forested catchments the NH₄ concentration level in freshwaters is generally low. The model slightly overestimated the runoff NH₄ concentrations and the NH₄ leaching with the present parameterisation (Fig. 5a and Table 6).

For the period 1995–1999, the mass balance for N was close to zero. INCA calculated a change of the N store of the catchment of 2.42 kg N ha⁻¹ over the whole period (Table 6). The measured NO₃ leaching of 16.8 kg ha⁻¹ for the five years (calculated with the discharge volumes adjusted to the Cl⁻ budget), was matched by the model.

The mean annual denitrification of 2.6 kg N ha⁻¹, calculated by INCA in the case of the Steinkreuz catchment with a given denitrification rate of 0.002 d⁻¹, corresponds to measurements in other beech forests (Borken and Brumme, 1997; Butterbach-Bahl *et al.*, 2001). Also the model parameterisation resulted in annual average plant uptake rates of 6.4 kg ha⁻¹ NO₃-N and of 2.4 kg ha⁻¹ NH₄-N, which corresponds to the total N demand for increment at older beech sites (Matzner, 1988).

In the case of the Lehstenbach catchment the temporal

Table 6. Simulated N fluxes

	Steinkreuz Σ 95–99 kg N ha ⁻¹	Lehstenbach Σ 95–99 kg N ha ⁻¹
Nitrate-N total	35.81	60.15
Ammonium-N total	43.11	46.29
Nitrate-N leaching	16.79	26.90
Observed values	16.79 ^a	24.64 ^b
Ammonium-N leaching	2.79	2.77
Observed values	0.32 ^a	1.12 ^b
Nitrate-N Uptake	32.06	19.29
Ammonium-N uptake	12.01	24.67
Ammonium-N nitrification	26.39	15.77
Nitrate-N denitrification	12.85	31.34
Δ Store	+2.42	+1.47

^a = corrected according to Cl⁻-budget

^b = 01/99-10/99: observed values; 11/99-12/99: estimated

pattern of discharge simulation corresponded in most cases to the observed changes (Fig. 5b). Deviations occurred in 1995 (between day number 145 and 230), where INCA failed to observe increases in discharge, and the simulated runoff exceeded the observed values in 1999.

The general concentration level for NO_3 in runoff is well matched. The measured annual average concentration of $\text{NO}_3\text{-N}$ was 0.92 mg l^{-1} while the model calculated 0.99 mg l^{-1} . Furthermore, the seasonality of the runoff NO_3 concentration, which is characterised by higher values in the winter and lower in summer, is reproduced well by the model.

As with the Steinkreuz, the model slightly overestimated the NH_4 concentrations in runoff and the NH_4 leaching with the present parameterisation. The model calculated a NO_3 flux with runoff of $26.9 \text{ kg N ha}^{-1}$ for the five years period, which is close to the observation (24.6 kg ha^{-1}).

For the period 1995–1999, INCA calculated a change of N store in the catchment of only $+1.47 \text{ kg N ha}^{-1}$ (Table 6) indicating that the mass balance of N was close to zero.

The model parameterisation resulted in a mean annual denitrification of 6.3 kg N ha^{-1} for the Lehstenbach. This reflects the obviously high N outputs by denitrification from the large wetland areas of the Lehstenbach catchment. High rates of denitrification were also postulated from comparing NO_3 in soil solution, groundwater and runoff in the Lehstenbach (Matzner *et al.*, 2001).

Given the present parameterisation, INCA calculated mean annual plant uptake rates of $3.9 \text{ kg ha}^{-1} \text{ NO}_3\text{-N}$ and of $4.9 \text{ kg ha}^{-1} \text{ NH}_4\text{-N}$. The overall rates of N uptake correspond to those estimated for comparable Norway spruce sites (Matzner, 1988).

Discussion

GENERAL

Both catchments studied receive high rates of N deposition as indicated by the observed N fluxes in throughfall which are typical for Central European conditions (Dise *et al.*, 1998a, b; Armbruster *et al.*, 2001). The N fluxes with throughfall underestimate the actual N input because of above ground uptake of N by needles and leaves. The extent of underestimation is difficult to assess (Harrison *et al.*, 2000). Beside deposition rates, both catchments have further similarities in their N turnover. The average annual NO_3 concentrations in runoff were between $0.7\text{--}1.4 \text{ mg l}^{-1}$ in both catchments, leading to average losses with runoff of 3.3 (Steinkreuz) to 4.9 (Lehstenbach) $\text{kg N ha}^{-1} \text{ yr}^{-1}$. Nitrate was the dominant N form in soil solutions and in runoff from both catchments. There was some seasonal variation

of NO_3 in runoff in both but no overall trend in NO_3 was found. In the case of the Lehstenbach, the lack of long term trend in runoff can be extended back to 1987 (Matzner *et al.*, 2001). Small seasonal variations of NO_3 concentrations in the runoff from forested catchments and the lack of long term trends were also reported from other European forested catchments (Wright *et al.*, 2002).

At the plot scale, however, the N turnover is significantly different at the spruce dominated Coulissenhieb plot compared with the beech dominated Steinkreuz plot. N retention was very low at the Coulissenhieb plot, indicating severe N saturation. Losses of N with seepage almost equalled N fluxes in throughfall. In contrast, the N retention at the Steinkreuz plot was much higher. There was no general difference in N retention between coniferous and deciduous tree species when a large number of plots and catchments in Europe were compared (Dise *et al.*, 1998a). The C/N ratios of the forest floor were rather low in both plots and, while not indicating differences in immobilization capacity of the soil, suggest a high risk of NO_3 leaching (Gundersen *et al.*, 1998). It is only possible to speculate on the retention mechanisms responsible for the N retention at the plot scale. Brumme *et al.* (1999) reported higher rates of N_2O emissions from beech sites compared to spruce sites, but the rates cannot account fully for the difference. The stand at the Coulissenhieb plot is older and less dense than that at the Steinkreuz plot which might result in higher plant uptake rates of N at Steinkreuz. However, the calculation of N retention is based on throughfall N fluxes which underestimate deposition rates. At the Coulissenhieb plot, the unknown N uptake by the canopy might be higher as compared to Steinkreuz because of the high fog frequency at these elevations (Klemm *et al.*, 2001).

The difference between N fluxes with seepage at the plot scale compared to runoff from the catchment were minor at Steinkreuz. At the Lehstenbach, however, the large difference between N fluxes with seepage from the terrestrial site and the N flux in runoff are influenced by the riparian zones which cover about 35% of the catchment. Denitrification is likely to play a significant role in regulating the N budget of the Lehstenbach catchment. There are no field measurements of denitrification available for the catchment studied here but a reduction in sulphate was reported in the riparian zones of the Lehstenbach catchment (Alewell and Giesemann, 1996; Alewell and Novak, 2001), indicating severe anoxic conditions in these areas.

Overall, the N budget of both catchments seems to be in a steady state where N sinks in the ecosystem (biomass and soil organic matter) and losses by denitrification match the present N input by deposition.

INCA APPLICATION

With a simplified parameter set, INCA matched the general level of NO_3 concentrations and fluxes in runoff at the Steinkreuz and the Lehstenbach catchments. The calculated annual loads for the single processes of the N-cycle, were realistic. The parameter settings resulted in a good representation of the seasonal patterns in NO_3 of the runoff from the Lehstenbach. However, the model was not able to follow the NO_3 concentrations in runoff at the Steinkreuz catchment during rewetting. This is an exceptional case since, in this very small catchment, the creek becomes dry in summer. The water fluxes at these particular periods were small and this had little effect on the annual runoff fluxes calculated by the model.

INCA estimated NH_4 leaching rates which are about $0.3\text{--}0.5 \text{ kg ha}^{-1} \text{ yr}^{-1}$ higher than those observed. The calibration of the model to realistic NH_4 leaching rates was difficult because the NH_4 plant uptake is zero in winter and nitrification is low in the model while leaching takes place during the whole year.

Papen and Butterbach-Bahl (1999) as well as Teepe *et al.* (2000) observed the highest N_2O emissions from forest soils during thawing periods in winter. This is not accounted for in the model since denitrification is inhibited in winter by low temperatures. However, this had no apparent influence on the model predictions of nitrate in runoff.

From the viewpoint of the N turnover in small homogeneous forested catchments, there are general shortcomings of the model. All parameters and rates are kept constant throughout the simulation and there is no dynamic feedback between inputs, environmental conditions and process rates. Those are, in fact, often not even known today in the case of forested catchments. For the parameterisation of the model, the rates of N transforming processes need to be known in advance or must be calibrated. Thus, most of the model results in terms of uptake, mineralisation, and denitrification and runoff losses are given as input and the predictive capability on N turnover and fluxes under changing conditions is limited.

The big advantage of the INCA model in application to forested catchments is that the prediction of NO_3 output is based on hydrological conditions allowing an analysis of NO_3 concentrations patterns in runoff throughout the year and under varying hydrological conditions. There is a clear and substantial use of the model in the calculation of critical loads of N deposition for forests. Models used presently are based on steady state mass balance using annual fluxes and without taking hydrological conditions into account (Werner *et al.*, 1999; van der Salm and de Vries, 2001). INCA seems to be an excellent tool to calculate NO_3

concentrations and NO_3 fluxes under a variety of deposition backgrounds and hydrological conditions with seasonal to daily resolution. This needs to be tested in future work.

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